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OESCHGER CENTRE CLIMATE CHANGE RESEARCH

Workshop on "Developments in isotope ratio measurements for gas analysis", Bern, 10.10.2019

Insights into the new tracer ¹⁷O-excess: results from the Swiss precipitation network

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Isotopologues of water: Tool to retrieve information about the water cycle



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Early steps in water isotope research R. Gonfiantini, H. Craig and L. I. Gordon





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Water cycle processes and its implication on isotopologues



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Isotopologues of water: Second order effects

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16

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d-excess in ice cores as tracer of

18

- Moisture source temperature (Ciais and Jouzel, 1994)
- Precipitation site temperature (Stenni et al., 2010, Vimeux et al., 2002)
- Relative humidity at evaporation site (Jouzel, 1982, Pfahl and Sodemann, 2014, Aemisegger and Soltje 2018)



 $\delta D - 8 \cdot \delta^{18}O = d$ -excess

(Dansgaard, 1964)

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Conclusion first part



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- Tools such as ¹⁷O-excess and d-excess are helpful in disentangling kinetic from equilibration processes.
- > We observed a clear anti-correlation between ¹⁷O-excess and the stable isotopes ($\delta^{17}O$, $\delta^{18}O$ and δD) in precipitation samples in contrast to
- > no or a slight positive correlation between ¹⁷O-excess and the stable isotopes $(\delta^{17}O, \delta^{18}O \text{ and } \delta D)$ in water vapor.
- The mean value of ¹⁷O-excess of the water vapor of about +20 per meg points to an important but rather weak (10 %) kinetic (diffusive) fractionation at the source location which might probably be present during the complete vapor path from the source location to the site of precipitation.
- The correspondent value for precipitation is about +45 per meg and is expected due to the higher slope of the equilibration fractionation of 0.529 between ¹⁷O and ¹⁸O.

$\delta \textbf{D}$ and d-excess values on Dome Fuji ice



Anti-correlation between δD and d-excess Could indicate an increased continentality effect!

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d-excess correlations with humidity and sea surface temperatures



FIG. 8. Correlation of (a) dc with SST and (b) dc with hs using monthly mean data and ak(MJ79) during SLOE. Both patterns look similar when using SLH-weighted monthly means. Land areas, regions with seasonal mean SLOE frequencies below 1% (tropics), or areas with seasonal sea ice are masked. \rightarrow proxy for relative humidity

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Water cycle processes and its implication on isotopologues



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Isotopologues of water: Second order effects

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¹⁷O-excess in vapour over the ocean is

18

- Anitcorrelated with the relative humidity (RH) (Uemura, 2010)
- Can it act as RH tracer?
- Is there more to say?



$^{17}\text{O-excess} = \ln (\delta^{17}\text{O} + 1) - 0.528 \cdot \ln (\delta^{18}\text{O} + 1)$

(Barkan and Luz, 2004)

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Primary and second-order effects



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$$1^{17}O-excess = -\ln({}^{18}\alpha_{eq}^{0.529}({}^{18}\alpha_{diff}^{0.518}(1-h_n)+h_n)) + 0.528 \cdot \ln({}^{18}\alpha_{eq}({}^{18}\alpha_{diff}(1-h)+h_n))$$

$$3004 \text{ E. Barkan and B. Luz}$$

From equation above: about -0.001 %/% = -1 permeg /%!!

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Demands on well-calibrated and precise instrumentations are high!!



Figure 1. The ¹⁷O-excess of atmospheric vapor in the ocean versus normalized humidity (h_n). Observed data (solid circles) with two outliers (open circles) are shown. The solid line is based on model calculations with optimized ¹⁸ α_{diff} (1.008). The chain line is based on model calculations with the generally accepted ¹⁸ α_{diff} values of 1.005.

Slope of -0.64permeg/%

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Slightly negative values for high humidities

Jungfraujoch monthly precipitation using a Picarro L2140-i instrument

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Mean of 47 ± 30 per meg \rightarrow 43 (+ 23, - 28) % humidity

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Jungfraujoch monthly precipitation and temperature

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Anticorrelation with temperature indirect effect

Jungfraujoch precipitation vs. temperature

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Jungfraujoch precipitation

Anticorrelation with temperature indirect effect



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Seasonality of Jungfraujoch precipitation

Seasonality with an amplitude of 34 per meg and a monthly uncertainty of ±3 per meg i.e. relative humidity changes 43 (+ 17, -14) %.

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Jungfraujoch monthly precipitation variations around the mean

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Decadal and interannual variations

Jungfraujoch monthly precipitation variations around the mean

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Dedacal trend (reason yet unkonwn)



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Jungfraujoch precipitation

Interannual variations (might be connected to NAO)



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Jungfraujoch precipitation

Intermonthly variations and measurement uncertainty

Swiss precipitation stations

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| Station | Abbreviation | Alt.m (a.s.l) | Long. WGS (deg) | Lat. WGS (deg) |
|--------------|----------------------|---------------|-----------------|----------------|
| Basel | BAS | 292 | 7.582 | 47.541 |
| Locarno | LOC | 379 | 8.787 | 46.172 |
| Bern | BER | 541 | 7.439 | 46.951 |
| Meirengen | MER | 598 | 8.178 | 46.727 |
| Guttannen | GUT | 1055 | 8.292 | 46.656 |
| Grimsel | GRI | 1980 | 8.332 | 46.571 |
| Jungfraujoch | $_{ m JFJ}$ | 3580 | 7.985 | 46.548 |

Swiss precipitation stations



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Seasonalities of Swiss precipitation stations

Amplitudes are between 28 and 52 per meg, means between 30 and 65 per meg.

Swiss precipitation stations



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Seasonality of Jungfraujoch precipitation

There might be constrasting dependencies at low and high altitudes, respectively (boundary layer influence, e. g. Kern et al, 2014)

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Difficulty of measurements, in particular at lower concentrations (vertical profile)

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Difficulty of measurements, in particular at low concentrations

Figure 1. Total uncertainty of δD , $\delta^{18}O$, and *d* excess over the range of H₂O_v mole fractions observed in flight.

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¹⁷O-excess is anticorrelated with $\delta^{18}O_{precip}$

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¹⁷O-excess is anticorrelated with $\delta^{18}O_{precip}$

Jungfraujoch trapped water vapor

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Thank you for your attention

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